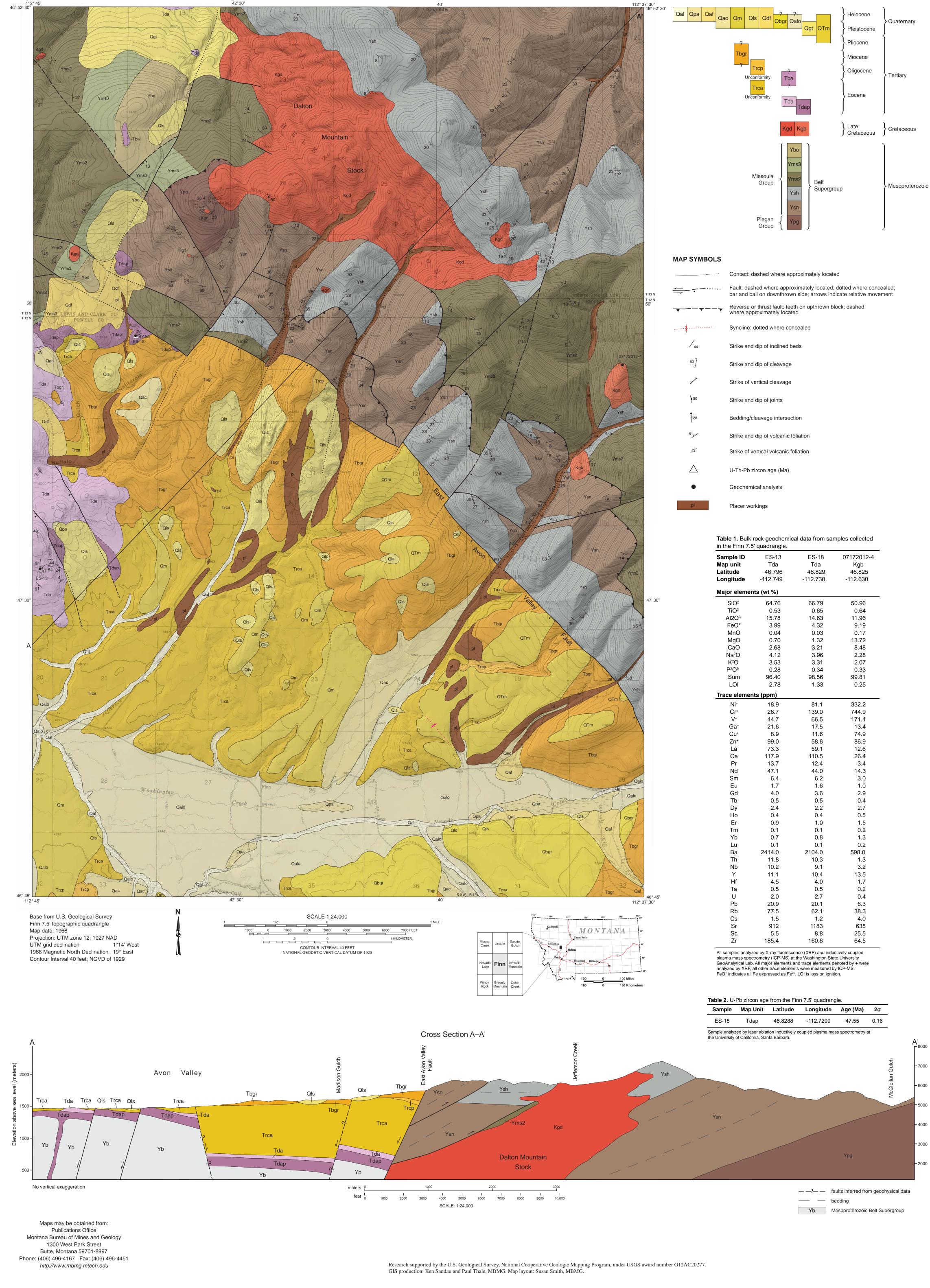
MONTANA BUREAU OF MINES AND GEOLOGY A Department of Montana Tech of The University of Montana



Map unit Latitude	Tda 46.796	Tda 46.829	Kgb 46.825
Longitude Major elemei	-112.749 nts (wt %)	-112.730	-112.630
SiO ²	64.76	66.79	50.96
TiO ²	0.53	0.65	0.64
AI2O ³	15.78	14.63	11.96
FeO*	3.99	4.32	9.19
MnO	0.04	0.03	0.17
MgO	0.70	1.32	13.72
CaO	2.68	3.21	8.48
Na ² O	4.12	3.96	2.28
K ² O	3.53	3.31	2.07
P^2O^5	0.28	0.34	0.33
Sum	96.40	98.56	99.81
LOI	2.78	1.33	0.25
Trace elemer	nts (ppm)		
Ni⁺	18.9	81.1	332.2
Cr⁺	26.7	139.0	744.9
V+	44.7	66.5	171.4
Ga⁺	21.6	17.5	13.4
Cu⁺	8.9	11.6	74.9
Zn⁺	99.0	58.6	86.9
La	73.3	59.1	12.6
Ce	117.9	110.5	26.4
Pr	13.7	12.4	3.4
Nd	47.1	44.0	14.3
Sm	6.4	6.2	3.0
Eu	1.7	1.6	1.0
Gd	4.0	3.6	2.9
Tb	0.5	0.5	0.4
Dy	2.4	2.2	2.7
Ho	0.4	0.4	0.5
Er	0.9	1.0	1.5
Tm	0.1	0.1	0.2
Yb	0.7	0.8	1.3
Lu Bo	0.1 2414.0	0.1	0.2
Ba Th	2414.0 11.8	2104.0 10.3	598.0 1.3
Nb	10.2	9.1	3.2
Y	10.2	10.4	
r Hf	4.5	4.0	13.5
Ta	4.5 0.5	4.0 0.5	0.2
Ia	0.0	0.0	0.2

CORRELATION DIAGRAM

INTRODUCTION

The Montana Bureau of Mines and Geology (MBMG), in conjunction with the STATEMAP advisory committee, selected the Finn 7.5' quadrangle in west-central Montana for detailed mapping to provide new information on stratigraphy and structure in an area where only reconnaissance mapping (Lewis, 1998; 1:250,000 scale) existed. The detailed mapping is part of the MBMG's effort to complete the Elliston 30' x 60' quadrangle (1:100,000 scale) geologic map. The northern, mountainous part of the Finn quadrangle is predominately U.S. Forest Service or Bureau of Land Management administered land; the southern part encompasses the north end of the Avon Valley (fig. 1) and

GEOLOGIC SUMMARY

is mostly privately owned.

The northern part of the Finn quadrangle is underlain by Mesoproterozoic Belt Supergroup metasedimentary rocks, including the middle Belt Piegan Group and the upper Belt Missoula Group. The Piegan Group, which includes the Helena and overlying Wallace Formations (Winston, 2007) occurs as poorly exposed calc-silicate rocks adjacent to the Dalton Mountain Stock. Because of the poor exposure and metamorphism, the individual formations could not be differentiated. Overlying the Piegan Group are the Snowslip, Shepard, Mountain Shields (members 2 and 3), and Bonner Formations, which are part of the upper Belt Missoula Group. The Missoula Group strata comprise an approximately 4,400 m section of quartzites, argillites, and dolomites.

The Belt strata were intruded by Late Cretaceous granodiorite of the Dalton Mountain Stock and smaller bodies of gabbro (table 1, map unit Kgb). These intrusive rocks are considered satellite bodies of the Cretaceous Boulder Batholith (Schmidt and others, 1994). Several gold mines and placer deposits are associated with the intrusive rocks.

The youngest rocks exposed in the quadrangle are Tertiary volcanics, and Tertiary and Quaternary valley-fill and surficial deposits. The volcanic rocks are predominantly aphanitic and porphyritic dacite (table 1, fig. 2), trachydacite, and minor basalt. A U-Pb zircon age from a porphyritic flow (map unit Tdap) in the NE¹/₄ of sec. 4, T. 12 N., R. 9. W., yielded an Eocene age of 47.55 ± 0.16 Ma (table 2). Tertiary valley-fill deposits include the Renova Formation (Eocene Climbing Arrow Member and Miocene/Oligocene Cabbage Patch member) and overlying boulder gravel deposits. Loen (1990) suggested a late Miocene (?) age for similar gravel deposits near Blackfoot City (south of the map area) and interpreted them as alluvial fans. These gravels have been mined for placer gold. Quaternary surficial deposits are widespread in the Avon Valley and include alluvium and colluvium along stream

drainages, landslides (primarily developed in the Renova Formation), debris flows, gravel mantle deposits, and glacial till in the northwest corner of the map. Structure

The Finn quadrangle is within the eastern part of the Lewis and Clark line (LCL)—a major western Montana WNW-striking tectonic zone recurrently active since the Proterozoic (Sears and Hendrix, 2004; Wallace and others, 1990; Reynolds, 1979). From Late Cretaceous to late Paleocene, the LCL underwent sinistral (left-lateral) transpression associated with crustal shortening (Sears and Hendrix, 2004). Beginning in late Eocene, deformation has been predominantly dextral (right-lateral) transtension associated with regional Basin and Range extension (Stickney, 2015; Sears and Hendrix, 2004). The major Cretaceous transpressive structures in the Finn quadrangle are northwest-trending thrust faults and regional-scale folds that deform the Belt rocks. The Late Cretaceous Dalton Mountain Stock intrudes the east limb of a broad regional syncline, is northwest-southeast elongate, and may have been intruded during folding and thrust faulting. Younger, NW- and NE-trending, high-angle transtensive oblique-slip faults cut the older structures.

The most prominent fault in the quadrangle, herein called the East Avon Valley Fault, is a down-to-the-southwest, range front fault that separates uplifted Belt rocks in the Nevada Mountains from Tertiary and Quaternary sediments in the Avon Valley. The fault is cut by steep NE-trending faults and forms the northeast boundary of the Avon Valley graben (Reynolds, 1979). Gravity data from the northern Avon Valley (Knatterud and others, 2015) suggest the northeast margin of the graben is a series of southwest-dipping normal or oblique-slip faults. Although NW-trending faults other than the East Avon Valley Fault were not recognized during mapping, faults inferred from the gravity data are shown on the cross section. The West Avon Valley Fault truncates the Miocene–Oligocene Cabbage Patch member of the Renova Formation and Tertiary alluvial fan deposits of probable Miocene age (Loen, 1990), constraining movement on the fault to Miocene or younger.

Although active faults have not been recognized in the Finn quadrangle, recent seismicity indicates continued tectonism on preexisting faults with oblique-normal slip (Stickney, 2015).

PREVIOUS MAPPING

The entire quadrangle was included on small-scale (1:250,000) maps by Lewis (1998) and Wallace and others (1987). Weber and Witkind (1979) completed a reconnaissance geologic map of the southwest quarter of the quadrangle that focused primarily on surficial deposits. Bierwagen (1964) and Melson (1964) mapped parts of the quadrangle's northeast corner.

DESCRIPTION OF MAP UNITS

- Qal Alluvium (Quaternary: Holocene)—Gravel, sand, silt, and clay along Nevada Creek and its tributaries. Clasts are generally rounded to subrounded, cobble size and smaller, but boulders also present. Thickness generally less than 10 m (33 ft).
- **Qpa Paludal deposit (Quaternary: Holocene)**—Argillaceous silt, sand, and organic matter deposited in pond or marsh environments associated with floodplain deposits of Nevada and Washington Creeks. Also associated with a landslide deposit south of Buffalo Creek. Thickness probably less than 5 m (15 ft). Alluvial fan (Quaternary: Holocene)—Gravel, sand, silt, and clay in fan-shaped deposits along Shingle Mill
- and Nevada Creeks in the southern part of the map. Large, locally derived clasts are both matrix- and clast-supported. Coarse clast size is dominantly cobbles, but also includes boulders and pebbles. Thickness generally less than 15 m (50 ft). Qac Alluvium and colluvium (Quaternary: Holocene)—Silt, sand, granules, and pebbles deposited on slopes by
- sheetwash alluvium incorporated with locally derived fine-grained colluvium. Deposits support hay fields along lower Washington Creek. Thickness generally less than 8 m (20 ft). Qm Mantle (Quaternary: Holocene and Pleistocene)—Deposits on pediments that include sheetwash alluvium,
- fine-grained colluvium, coarse-grained (pebble, cobble, small boulder) lag from older debris-flow deposits, and regolith. Thickness generally less than 6 m (20 ft).
- Qls Landslide (Quaternary: Holocene or Pleistocene)—Mass-wasting deposits of rotated or chaotic beds. The bentonitic Climbing Arrow Member of the Renova Formation (Trca) and units that immediately overlie it are particularly prone to developing landslide deposits. Thickness generally less than 60 m (200 ft). Qdf Debris-flow (Quaternary: Pleistocene or Holocene)—Matrix-supported deposits of local, upslope-derived
- volcanic rock and sparse sedimentary rock. Clasts range from granules to boulders. Thickness generally less than 20 m (75 ft). **Boulder gravel (Quaternary: Holocene(?) and Pleistocene**)—Unconsolidated deposits along Nevada Creek
- of reworked Tbgr. Dominantly weakly stratified rounded pebbles and cobbles in a coarse-grained matrix. Deposits southwest of Dalton Mountain include abundant subangular and angular Belt quartzite boulders as lag on the forest floor, but exposures were not observed. Thickness less than 20 m (70 ft).
- Qalo Alluvium, older than Qal (Quaternary: Pleistocene)—Fluvial deposits in the Avon Valley in channels wider than those occupied by the present-day underfit streams. Clasts are generally rounded to subrounded cobbles and smaller, but boulders also present. Thickness not known.
- **Q**gt **Glacial deposit (Quaternary, Pleistocene)**—Light brown, brown, and reddish brown granules, pebbles, cobbles, and boulders in a very sandy to silty matrix which locally is clayey and compact. Dominant lithologies are quartzite and argillite. Confined to Willow Creek drainage in northwest corner of map. **QTm** Mantle (Quaternary and Tertiary)—Bouldery deposits on slopes near the East Avon Valley Fault on the
- east side of the Avon Valley. Upslope bedrock on the footwall block is not necessarily the source of the deposits.

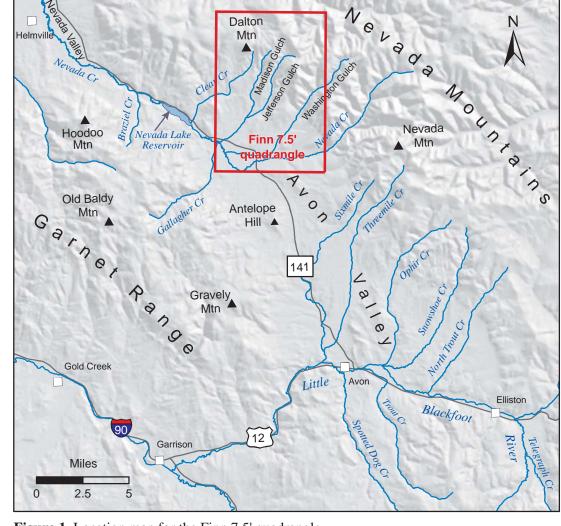


Figure 1. Location map for the Finn 7.5' quadrangle.

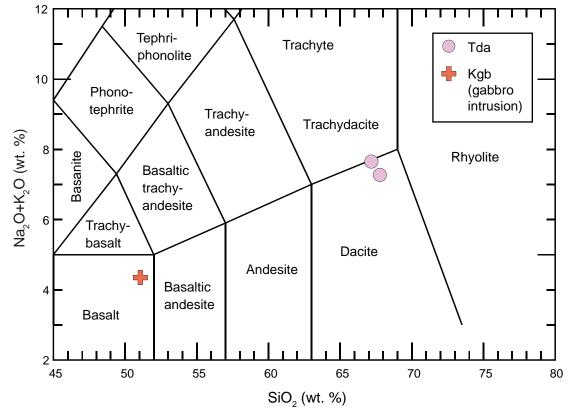


Figure 2. Whole-rock geochemical data from samples collected in the Finn 7.5' quadrangle. Data are plotted on a total alkali silica diagram (Na₂O + K_2O vs. SiO₂) after Le Bas and others (1986).

Tbgr	Boulder gravel (Tertiary) —Dominantly matrix-supported, b occur adjacent to East Avon Valley Fault. Clasts range from p matrix. Clast composition dominantly Belt Supergroup quartz Matrix in many places has incorporated sediment from the Cl (Trca). Deposit has been interpreted as Miocene (Loen, 1990; (Loen, 1990). May be as thick as 300 m (985 ft).
Trcp	Renova Formation, Cabbage Patch member, informal (Ter tuffaceous, micaceous, siltstone and fine-grained sandstone w slumped downslope along the west side of Jefferson Creek. It along Cow Creek where it contains numerous fossil root trace Gravely Mountain quadrangle to the south were mapped (Mos of the Renova Formation (Trcp). Thickness probably less than
Trca	Renova Formation, Climbing Arrow Member (Tertiary: E olive brown, and medium light gray bentonitic mudstone, yell gray sandstone and gravel/conglomerate lenses. Sandstone do quartz, feldspar, biotite, and rhyolite grains. Gravel/conglome rounded granules, pebbles, and cobbles of rhyolite porphyry. A occurs a short distance above the mouth of Jefferson Creek (P as a coalbed near Finn mined for domestic use (Glenna Stucko of Jefferson Creek, fish scales occur in platy siltstone and pap Bentonitic mudstone typically displays "popcorn" weathering landslide development. Springs in NW ¹ /4 sec. 9, and NW ¹ /4 sec crusts of Climbing Arrow Member. Unit referred to as "Black exposed. Exposed thickness approximately 180 m (600 ft).
Tba	Basalt (Tertiary: Eocene) —Black, fine-grained, massive bas phenocrysts. Locally vesicular. Weathers to light brown, angu equivalent to basaltic-andesite flows in the Gravely Mountain zircon age of 30.00 ± 0.19 Ma (Mosolf, 2015) and basalt in th others, 2016) with a K-Ar (whole rock) age of 32.3 Ma (Mitch
Tda	Dacite (Tertiary: Eocene) —Gray and dark gray weathering, trachydacite lava flow with autobrecciated flow bottoms over vesiculation are rarely preserved. Individual flows commonly flaggy, angular talus, often with red iron staining on parting su trachytic texture consisting mainly of plagioclase microlites, b Slightly porphyritic flows exhibit a similar trachytic texture are subhedral to euhedral phenocrysts of plagioclase, hornblende,
Tdap	Dacite, porphyry (Tertiary: Eocene) —Gray, green, and red (table 1, fig. 2) with a distinct coarse, porphyritic texture. Law phenocrysts of plagioclase (up to ~5 mm) and minor amounts groundmass commonly has a trachytic texture and consists ma also includes amphibole, biotite, and magnetite. Carapaces of encompass coherent interiors of individual lava flows. This up outcrops forming hoodoos and spires. A sample from the NE ¹ / ₂ zircon age of 47.55 \pm 0.16 Ma (table 2).
Kgd	Granodiorite (Late Cretaceous) —Medium to dark gray, equ of plagioclase, potassium feldspar, quartz, biotite and hornble Stock and associated smaller intrusions. Schmidt and others (Ma.
Kgb	Gabbro (Late Cretaceous) —Dark gray to greenish black, eq coarse-grained quartz gabbro (table 1, fig. 2) composed prima biotite and magnetite. Occurs as two small bodies along the W
Ybo	Bonner Formation (Mesoproterozoic) —Pink to red, medium quartzite. Abundant trough and planar crossbeds in beds typic separated by thin (0.1–1.0 cm), red to maroon argillite beds. O pebbles. A slabbed and stained sample contained 55 percent q percent plagioclase. Upper contact not exposed but is about 55 Nevada Lake quadrangle (McDonald and others, 2016).
Yms3	Mount Shields Formation, member 3 (Mesoproterozoic)— quartzite to argillite couples and couplets with abundant desic salt casts. Includes green interbeds and some red microlamina
Yms2	Mount Shields Formation, member 2 (Mesoproterozoic)— coarse-grained, feldspathic quartzite. Abundant planar and tro granules. Lower part is thinner bedded, consisting of couples thin red argillite. Bases of quartzite beds contain abundant red top part of the Mount Shields Formation, member 1 of Schmis samples from the Ophir Creek 7.5' quadrangle east of the Avo percent potassium feldspar, and 20 percent plagioclase (Lonn ft).
Ysh	Shepard Formation (Mesoproterozoic) —Dolomitic and nor argillite in microlaminae. Couplets of non-dolomitic red quart thin plates that, when dolomitic, have a characteristic orange- are common; mudcracks are rare. The upper half of the format fine-grained feldspathic quartzite originally included in Moun others (1994). However, we place the formation's upper conta distinctive rose-colored, dolomitic siltite–argillite couplets as

Mountain Stock. Thickness about 610 m (2,000 ft). Snowslip Formation (Mesoproterozoic)—Interbedded intervals of quartzite to red argillite couplets, dark green siltite to light green argillite couplets, and microlaminated argillite. Desiccation cracks and mud rip-up clasts are common. Argillite beds often contain irregular "bumps" thought to be ill-defined salt casts or related to microbial mats. Contains beds and lenses of distinctive white, coarse-grained, well-sorted, feldspar-poor quartzite containing some well-rounded, frosted quartz grains. Lower 50 m (160 ft) dominated by microlaminated green dolomitic siltite and argillite. Upper 50–75 m (160–245 ft) is red, flat laminated, medium-grained quartzite in beds 0.5–1.0 m thick. Thickness approximately 1,430 m (4,700 ft).

are contact metamorphosed to pale to dark green hornfels with diopside and/or tremolite. Thickness not determined.

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MBMG Counties, Montana and Susan M. Vuke

MBMG Open-File Report 672 Geologic Map of the Finn 7.5' Quadrangle, Lewis and Clark and Powell Catherine McDonald, Jesse G. Mosolf,

2016

MBMG Open-File Report 672; Plate 1 of 1 Geologic Map of the Finn 7.5' Quadrangle, 2016

, but locally clast-supported deposits that primarily pebbles to large boulders in a clay/silt/sand rtzite with subordinate volcanic and plutonic rock. limbing Arrow Member of the Renova Formation); Schmidt and others, 1994). Contains placer gold

ertiary: Miocene and Oligocene)—Light tan, vith local lenses of rip-up-clast breccia. Unit has It is also present in southwest corner of map and aces. Similar deposits on Antelope Hill in the Iosolf, 2015) as part of the Cabbage Patch member an 45 m (150 ft) in map area.

Eocene)—Pale olive, light olive gray, moderate ellowish gray siltstone, and subordinate brownish lominantly medium- to coarse-grained, with nerate consists of clast-supported lenses of A coalbed approximately 0.3 m (1 ft) thick (Pardee and Schrader, 1933) and may be the same ckey, oral commun., 2012). West of Finn and east aper shale beds in a road cut along Highway 141. ng, large desiccation cracks, and is prone to sec. 22, T. 12 N., R. 9 W. are concealed by thin ckfoot City beds" by Loen (1990). Base of unit not

asalt flows with scattered olivine and pyroxene gular to sub-rounded blocks. Perhaps age in quadrangle to the south that yielded a U-Pb the Nevada Mountain quadrangle (McDonald and tchell Reynolds, written commun., 2014).

aphanitic to slightly porphyritic dacite to rlain by massive coherent flow interiors; zones of y exhibit distinctive flow banding and form surfaces. Aphanitic lavas exhibit a strong , but also pyroxene, magnetite, and volcanic glass. and groundmass mineralogy, but contain , and some biotite.

weathering dacite to trachydacite lava flows ava flows contain subhedral to euhedral ts of amphibole, biotite, and quartz. The aphanitic mainly of aligned microlites of plagioclase, but autobreccia several meters thick commonly unit weathers into blocks or plates, with some E¹/4 sec. 4, T. 12 N., R. 9. W., yielded a U-Pb

uigranular to porphyritic granodiorite composed ende; forms main body of the Dalton Mountain (1994) report a K-Ar age on biotite of 77.9 ± 2.8

quigranular to porphyritic, medium- to narily of pyroxene and plagioclase with minor Washington Creek drainage.

ium- to coarse-grained, poorly sorted, feldspathic cally 0.5–1.0 m thick. Quartzite beds often Contains sparse subangular granules and small quartz, 35 percent potassium feldspar, and 10 550 m (1,800 ft) thick in the west-adjacent

-Grayish red to blackish red, fine-grained iccation cracks, mudchips, and diagnostic cubic nae. Thickness about 275 m (900 ft).

–Pink to red, poorly sorted, medium- to ough crossbeds. Contains sparse subangular of white to pink, medium-grained quartzite and ed mudchips. The lower part was included in the nidt and others (1994). Two slabbed and stained von Valley contained 65–75 percent quartz, 5–15 and Vuke, 2015). Thickness about 520 m (1,700

on-dolomitic, dark green siltite and light green rtzite to argillite. Poorly exposed, but weathers to e-brown weathering rind. Ripples and load casts nation contains intervals of pink to gray, unt Shields Formation, member 1 by Schmidt and ntact at the top of a 50-m (160-ft) thick interval of distinctive rose-colored, dolomitic siltite-argillite couplets as was done in areas to the west (Lonn and others, 2010). The lower contact is placed at the bottom of the lowest dolomite-bearing beds. A stromatolite bed is often found near the base. Metamorphosed to pale green, striped hornfels near the contact with the Dalton

Piegan Group (Mesoproterozoic)—Adjacent to the Dalton Mountain Stock, carbonates of the Piegan Group